



## The dependence of Aerosol-Cloud Indirect Effects on the representation of the autoconversion: Formulation and sensitivity experiments in CESM

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# **Aerosol Effects on Clouds**



Polluted air (many CCN)



### Aerosol – Cloud – Interactions (ACI) Smaller drops => brighter clouds (S.Twomey 1977) => delay in precipitation (B.Albrecht, 1989)

# **Climate Forcing**



#### Greenhouse gases 3.0 ± 0.8 Wm<sup>-2</sup>

Aerosols -0.8 ± 1.2 Wm<sup>-2</sup>

IPCC, 2013, SPM.5

# Community Earth System Model (CESM)

- CESM is a fully-coupled, global climate model that provides state-of-the-art simulations of the Earth's climate (developed at NCAR)
- explicitly simulates Aerosol-Cloud Interactions making it possible to simulate aerosol indirect effects



# **Aerosol Cloud Interactions in CESM2**



I. Activation (CCN) = f(RH,w) W at cloud scale is critical

- 2. Autoconversion (loss process) is a function of N<sub>c</sub> (=ACI)
- **3. Accretion depends on q**<sub>r</sub>

## **Autoconversion parameterization**

Khairoutdinov and Kogan scheme (KK2000)

$$\boldsymbol{A} = \boldsymbol{k} \boldsymbol{q}_{c}^{a} \boldsymbol{N}_{c}^{-b}$$

A = autoconversion rate q<sub>c</sub>= cloud liquid (kg/kg) N<sub>c</sub>= droplet number (#/kg)





# **Obs: Holuhraun eruption, Iceland (2014-2015)**



## Holuhraun eruption: Model versus obs anomalies



Malavelle et al 2017, Nature

**CESMI** overestimates the change in LWP

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#### Sensitivity experiments: Varying exponent b

b	
1.79	KK2000
1.1	Wood, personal communication
0.5	Extreme value

+ adjust k to produce the same autoconversion rate

# **Sensitivity Experiments**

$$\boldsymbol{A} = \boldsymbol{k} \boldsymbol{q}_{c}^{a} \boldsymbol{N}_{c}^{-b}$$

5-year runs with prescribed climatological present-day SSTs Similar climate after retuning (k and Dcs)

exp b	k/1350	SWCF W/m2	LWCF W/m2	LWP	Dcs microns
0.5	0.002	-47.4	24.1	68.I	540
1.1	0.01	-48.4	24.8	68.4	540
1.79	0.06	-48.7	24.3	66.4	300

### **SWCF** bias compared to **CERES-EBAF**: Similar bias







80 60 50 40 30 20 10 -10 -20 -30 -30 -50 -60 -80

## Estimation of the aerosol indirect effect

Compare two simulations < present day aerosol (2000) pre-industrial aerosol (1850)

**Aerosol Optical Depth** 

2000

1850





2000 - 1850

Look difference between 2 runs: ΔRESTOM => Aerosol total effect ΔSWCF => Cloud albedo effect (1<sup>st</sup> indirect effect) ΔLWP => Cloud lifetime effect (2<sup>nd</sup> indirect effect)

## Estimation of the aerosol indirect effect

present day aerosol (2000) Compare two simulations < pre-industrial aerosol (1850)

**Aerosol Optical Depth** 

1850



	b = 1.79	b = 1.1	b = 0.5	
∆RESTOM (W/m2)	-1.56	-1.27	-1.18	Total effect
∆SWCF (W/m2)	-1.29	-1.17	-1.11	l <sup>st</sup> indirect effect
Δ <b>LWP (%)</b>	7.3%	4.72%	2.35%	2 <sup>nd</sup> indirect effect

### Difference in SWCF and LWP in 2000-1850 aerosol



### Impact on 20<sup>th</sup> century surface temperature

Has a large impact on the period 1940-1960



### 20<sup>th</sup> century variability

#### Impact is not within the variability range



## **Summary**

Malvelle, 2017: the Holuhraun eruption reduced the size of liquid cloud droplets had no discernible effect on other cloud properties like LWP.

Autoconversion KK 2000:  $A = k q_c^a N_c^{-b}$ 

Sensitivity tests with b = 0.5, 1.1, 1.79 impacts indirect effects (change in SWCF and LWP)

Change in b has a direct impact on the period 1940-60